

# Engineering Reconnaissance following the 2016 M 6.0 Central Italy Earthquake: Ver 1

**DOI:10.18118/G65K5W**

**GEER Team Leader: Jonathan P Stewart, University of California Los Angeles (UCLA)**

**Co-Team Leader: Giuseppe Lanzo, Università degli studi di Roma *La Sapienza***

## **Team Members:**

**Stefano Aversa, Università degli studi di Napoli *Parthenope***

**Francesca Bozzoni, Eucentre, Pavia**

**Filiberto Chiabrando, Nives Grasso, Vincenzo Di Pietra, Politecnico di Torino**

**Shideh Dashti, University of Colorado Boulder**

**Luigi Di Sarno, Maria Giovanna Durante, Università degli studi del Sannio, Reluis Consortium**

**Sebastiano Foti, Politecnico di Torino**

**Kevin Franke, Brandon Reimschiessel, Brigham Young University**

**Fabrizio Galadini, Emanuela Falcucci, Stefano Gori, Istituto Nazionale di Geofisica e Vulcanologia (INGV)**

**Robert E Kayen, US Geological Survey (USGS)**

**George Mylonakis, Elpida Katsiveli, University of Bristol, University of Patras**

**Alessandro Pagliaroli, Silvia Giallini, Università degli studi di Chieti-Pescara**

**Giuseppe Scasserra, Filippo Santucci de Magistris, Massimina Castiglia, Università degli studi del Molise**

**Stefania Sica, Michele Mucciacciaro, Università degli studi del Sannio**

**Francesco Silvestri, Anna D'Onofrio, Anna Chiaradonna, Filomena De Silva, Università degli studi di Napoli *Federico II***

**Armando L Simonelli, Augusto Penna, Università degli studi del Sannio, Reluis Consortium**

**Paolo Tommasi, Consiglio Nazionale delle Ricerche**

**Paolo Zimmaro, University of California Los Angeles (UCLA)**

**Geotechnical Extreme Events Reconnaissance Association**

**Report No. GEER-050A**

**15 September 2016**

## Introduction

The central Italy earthquake occurred on 24 August 2016 at 03:36 AM local time. The magnitude is listed as **M6.0** by [INGV](#) and **M6.2** by [USGS](#). Although initially reported as occurring at relatively shallow depths, the current source model from INGV places the hypocentral depth at 8 km, which is not especially shallow for shallow crustal earthquakes.

The earthquake was located in a gap between two earlier damaging events, the 1997 **M6.1** Umbria-Marche earthquake to the north-west and the 2009 **M6.1** L'Aquila earthquake to the south-east. This gap had been recognized prior to the event as a zone of elevated risk (GdL Istituto Nazionale di Geofisica e Vulcanologia, hereafter INGV, 2016). The present event and those that preceded it occurred along the spine of the Apennine Mountain range on normal faults and had rake angles ranging from -80 to -100. Each of these events produced substantial damage to local towns and villages; the present event most strongly affected Arquata del Tronto, Accumoli, Amatrice, and Pescara del Tronto, with a loss of life as of this writing of 294, generally from collapses of unreinforced masonry dwellings.

The NSF-funded Geotechnical Extreme Events Reconnaissance (GEER) association, with co-funding from the B. John Garrick Institute for the Risk Sciences at UCLA and the NSF I/UCRC Center for Unmanned Aircraft Systems (C-UAS) at BYU, mobilized a US-based team to the area from 5-9 September 2016. The US team worked in close collaboration with Italian researchers organized under the auspices of the Italian Geotechnical Society, the [Italian Center for Seismic Microzonation and its Applications](#), the Consortium ReLUIS, Centre of Competence of Department of Civil Protection and the Disaster RECOVERY Team of Politecnico di Torino. The objective of the Italy-US GEER team was to collect and document perishable data that is essential to advance knowledge of earthquake effects, which ultimately leads to improved procedures for characterization and mitigation of seismic risk.

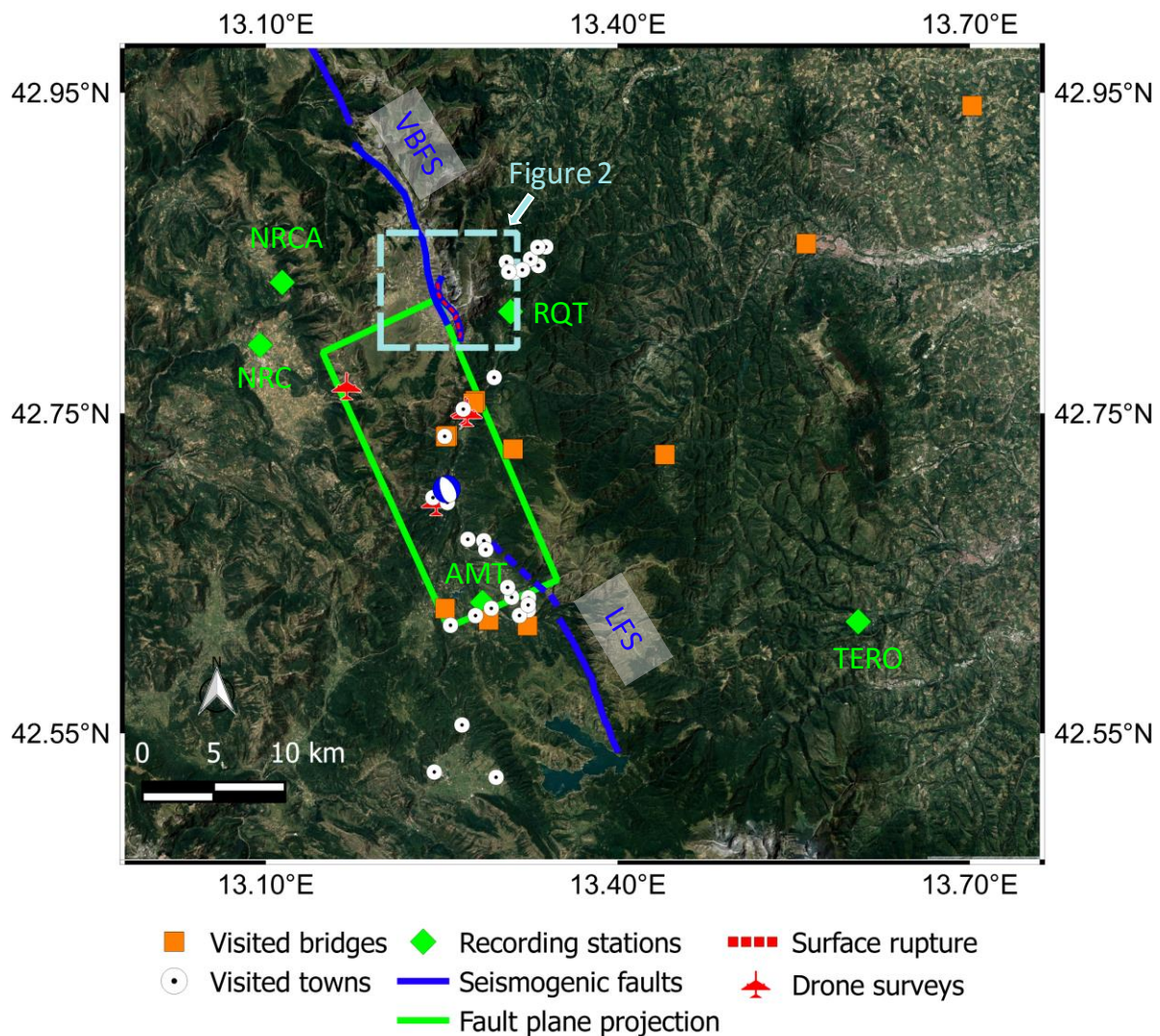
The Italy-US GEER team was multi-disciplinary, with expertise in geology, seismology, geomatics, geotechnical engineering, and structural engineering. Our approach was to combine traditional reconnaissance activities of on-ground recording and mapping of field conditions, with advanced imaging and damage detection routines enabled by state-of-the-art geomatics technology. This combination of reconnaissance techniques provides opportunities for innovative future study.

The objective of this brief report is to provide to the technical community, emergency responders, and public an account of our activities and preliminary findings in a timely way. A more complete presentation of significant aspects of this event and our detailed findings will be presented in a subsequent Version 2 report.

## Executive Summary of GEER Team Activities and Preliminary Findings

Figure 1 shows the most strongly affected region. Our activities focused on the following aspects of the earthquake event:

1. Surface fault rupture
2. Recorded ground motions
3. Locations of landslides and rockfalls. Mapping specific case histories.
4. Performance of bridge structures
5. Performance of building structures, with an emphasis on damage patterns

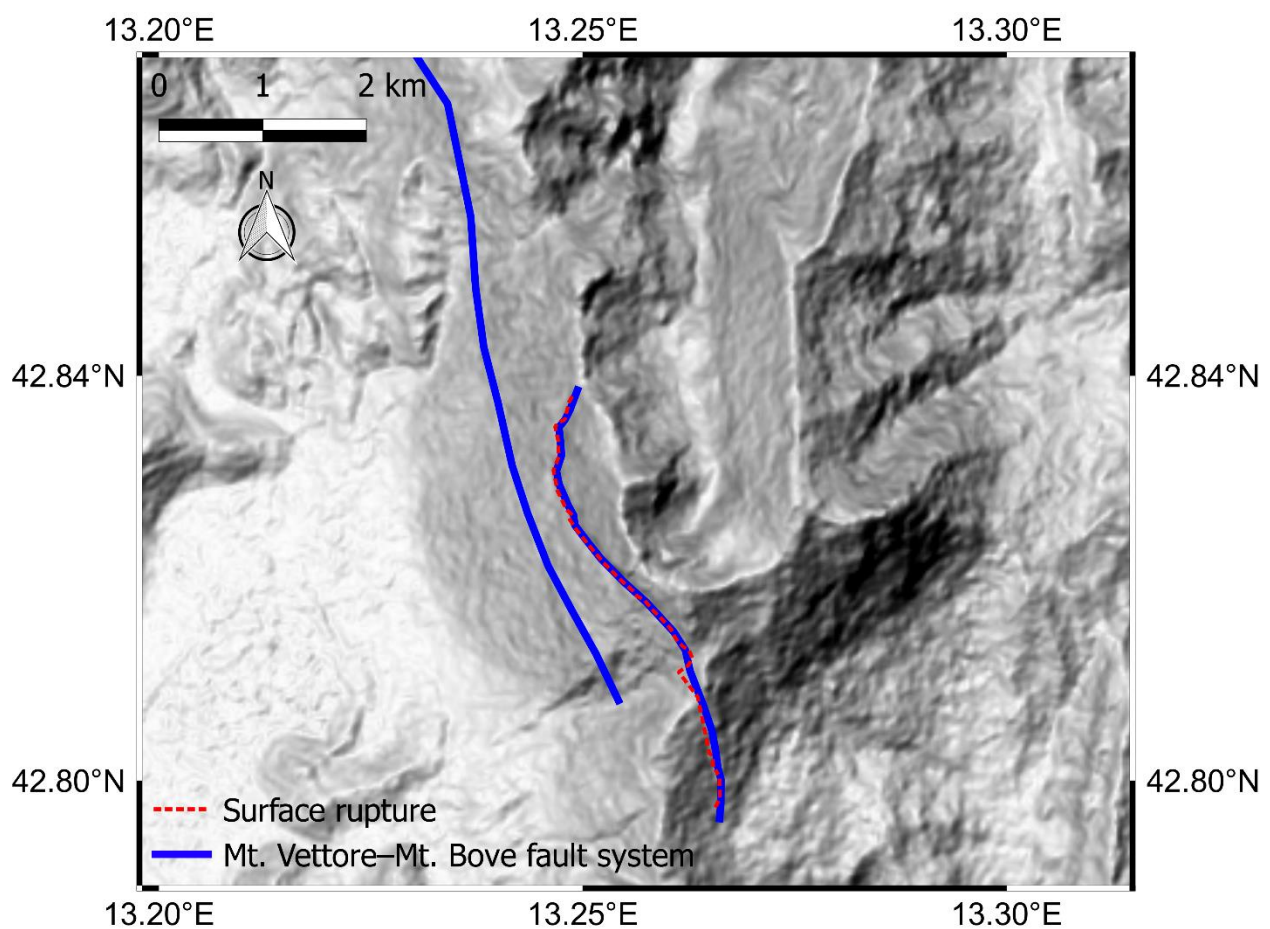


**Figure 1.** Regional map showing the Mt. Vettore–Mt. Bove fault system (VBFS, modified from Pizzi and Galadini, 2009), and the Laga fault system (LFS, modified from Galadini and Galli, 2003), INGV finite fault model and epicenter (GdL INGV, 2016), ground motion station locations, and locations of various earthquake effects discussed in this report. (Figure prepared by: P Zimmaro, E Falcucci, F Galadini, and S Gori).

## Surface Faulting

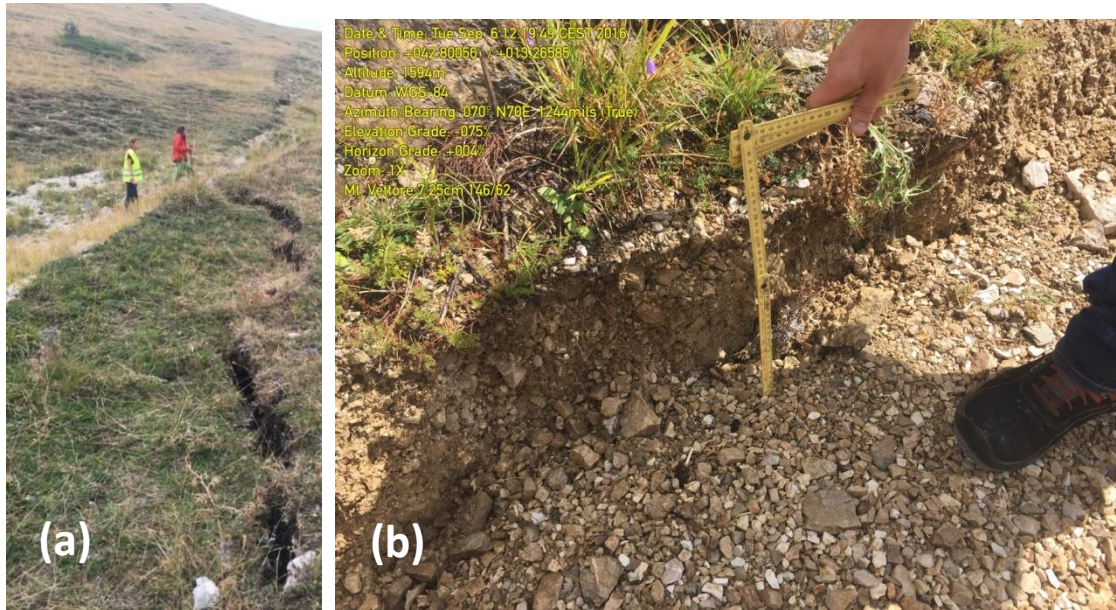
The mainshock occurred in either a bend or a stepover zone between two previously mapped normal faults – the Mt. Vettore fault and the Amatrice fault, which is the north extension of the Laga Mountains fault. These faults are shown in Figure 1.

Surface rupture occurred over a 4.8 km portion of the Mt. Vettore fault, as shown in Figure 1 and 2. Working with researchers at INGV (especially co-authors Galadini, Gori, and Falcucci), displacements on the primary rupture were measured as generally in the range of 10-25 cm. These displacements were down-dip, generally with no appreciable along-strike component. Figure 3 shows a general view of the fault trace (Figure 3a) and a typical displacement measurement on the fault (Figure 3b). We did not see secondary breaks on the hanging wall. Aerial imagery of the surface rupture is pending.



**Figure 2.** Detail map of surface fault rupture and pre-event mapping of Mt. Vettore-Mt. Bove fault system (adapted from Pizzi and Galadini, 2009). (Figure prepared by: P Zimmaro, E Falcucci, F Galadini, and S Gori).





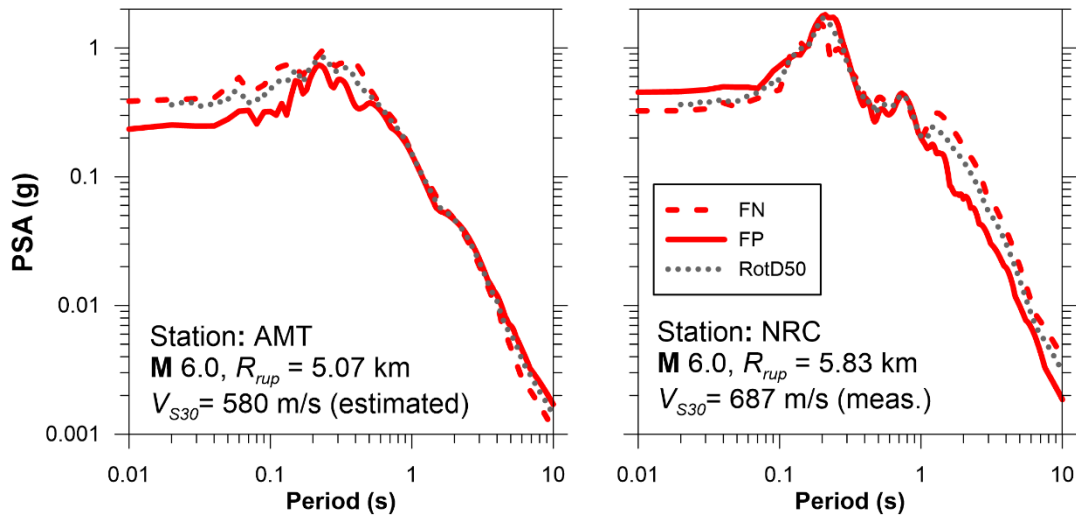
**Figure 3.** (a) photographs of surface rupture on trace of Mt. Vettore fault (N42.79953, E13.26634, 5 September 2016) and (b) typical measurement (N42.8056, E13.26585, 6 September 2016).

### Ground Motions

Two ground motion networks operate widely in Italy: INGV and Dipartimento della Protezione Civile (DPC). Data from these networks are disseminated at <http://esm.mi.ingv.it>.

In collaboration with the Pacific Earthquake Engineering Research (PEER) center, we downloaded volume 1 (digital uncorrected) data from the mainshock and aftershocks on 8/24 and 8/26 with INGV magnitudes **M**5.4, and 4.8. In total, 257 records were downloaded on 1 September 2016 and processed using standard PEER processing (Ancheta et al. 2014), which removes any static offset effects that might otherwise be present. We are aware of 105 additional records now available, which will be considered subsequently. We are presently compiling site parameters for stations without shear-wave velocity measurements, which may include some geophysical testing using surface wave methods at selected sites. Our Ver 2 report will present analysis of the data.

Figure 4 shows pseudo-acceleration response spectra (PSA) from the two stations most proximate to the mainshock rupture plane (AMT and NRC), which are shown in Figure 1. The corrected ground motions have been rotated into fault normal (FN) and fault parallel (FP) components. The two ground motions are of comparable amplitudes (NRC slightly higher), despite substantially different damage levels (details below). The AMT ground motion shows evidence of polarization in the FN direction at short oscillator periods (< 1.0 sec), while the NRC motion is stronger in the FN direction at long periods (> 1.0 sec).



**Figure 4.** Mainshock pseudo acceleration response spectra (5% damping) for Amatrice (AMT) and Norcia (NRC) sites. Fault strike taken as 150 deg for computation.

We have visited the AMT and NRC sites, and further site visits to accelerographs are likely in subsequent phases of work. Generally, the damage patterns near the AMT site were also suggestive of stronger shaking in the FN direction, which will be detailed in our Ver 2 report. We found no particularly noteworthy attributes of the recording shelters at these sites.

### Landslides and Rockfalls

Figure 5 shows locations of known rockfalls and landslides. Prior to deployment of the GEER team, we reviewed reports of rockfalls from ISPRA (2016) and CERI (2016). We found a limited number of landslides beyond those identified in the preliminary reports.

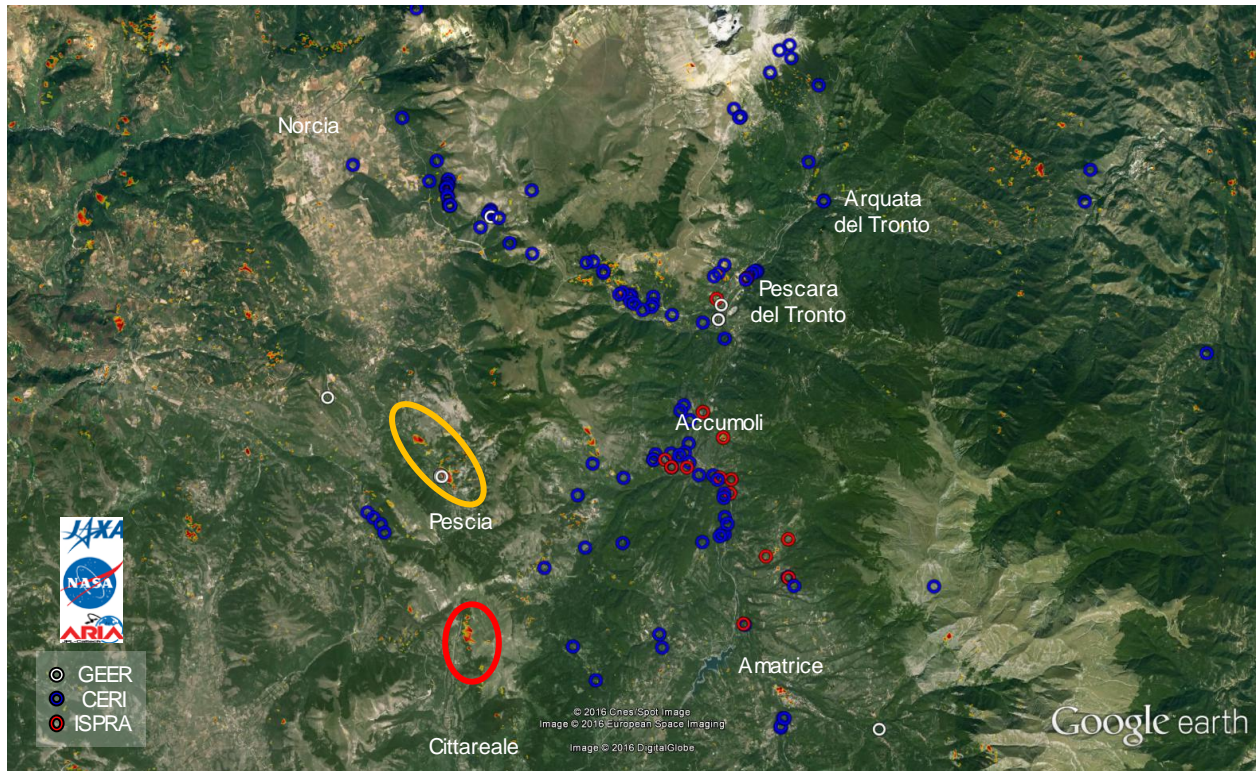
We overlaid the ISPRA and CERI locations on a map of multi-epoch (pre- to post-earthquake) deformation based on InSAR coherence changes (Damage proxy maps, ARIA project, JPL-Caltech, available at: <http://aria-share.jpl.nasa.gov/>), as shown in Figure 5. Our approach to reconnaissance activities was to initially visit sites of mapped deformation, but no prior rockfall observations, followed by more detailed mapping activities for high-value case histories.

Using direct visual observation where possible, in combination with aerial drone based imaging, we found no detectable landslide activity in the high-deformation zones shown in Figure 5 without prior landslides.

Using a DJI™ Phantom 4 drone, a customized Align™ T-Rex 800 drone, and an Ebee Sense Fly drone, we imaged the four locations shown in Figure 1. These data will be used also for for high resolution digital map creation (orthophoto) Digital Surface Model (DSM) and 3D models of the acquired areas. Information on where to find down-sampled video for two of these locations (Accumoli and Pescara del Tronto) are provided on the Central Italy Earthquake event page within the GEER Association website ([www.geerassociation.org](http://www.geerassociation.org)). Figure 6 shows aerial photos taken by drones at Accumoli (Figure 6a) and Pescara del Tronto (Figure 6b).



Observed landslides were mainly small rock failures (wedge slides, topples and slides in intensely fractured/weak rocks) which very often generated rockfalls. One of the largest failures occurred on a slope that, in the past, was involved in similar instability phenomena (Pescara del Tronto, just above the Salaria State road, Figure 6b). In several instances (specifically Tufo, Pescara del Tronto and Accumoli, Figure 6a), retaining structures underwent displacements and rotation (toppling) with consequences to the backfill and hence to the areas behind (including roads).



**Figure 5.** Mapped rockfalls and landslides from ISPRA (red circles), CERl (blue circles), and GEER (this study; white circles), along with the damage proxy map of the area produced by the ARIA project (JPL-Caltech).





**Figure 6.** Aerial photos taken by drones at: (a) Accumoli (N42.69425, E13.24993, 9 September 2016); and (b) Pescara del Tronto (N42.750401, E13.272109, 9 September 2016).

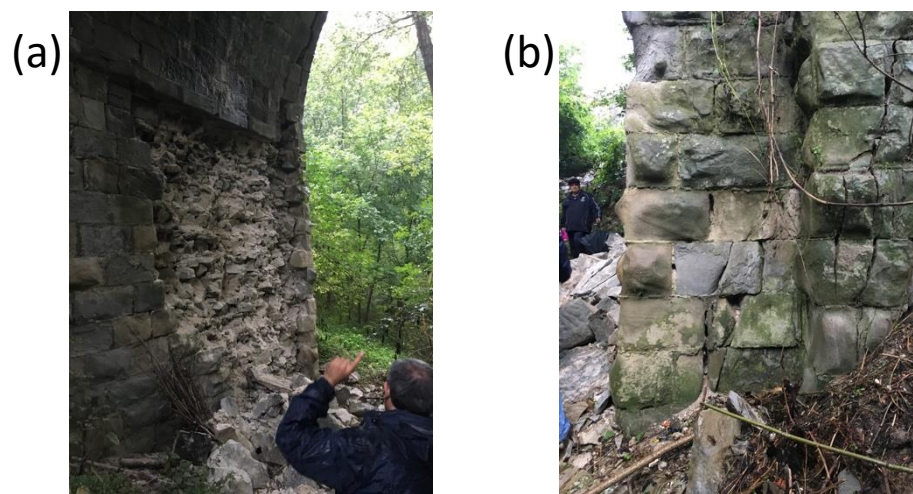


### Bridges (collaboration with [Reluis](#), responsible investigator: L Di Sarno)

GEER worked in close collaboration with Consortium ReLUIS to inspect 11 bridges at the locations shown in Figure 1. These bridges have a variety of configurations, but can be broadly viewed as relatively contemporary structures (constructed since approximately the 1960s) and older arch masonry structures. The contemporary bridges inspected are reinforced concrete (RC) and composite RC and steel structures, generally built in the 1960s. The contemporary bridges inspected include:

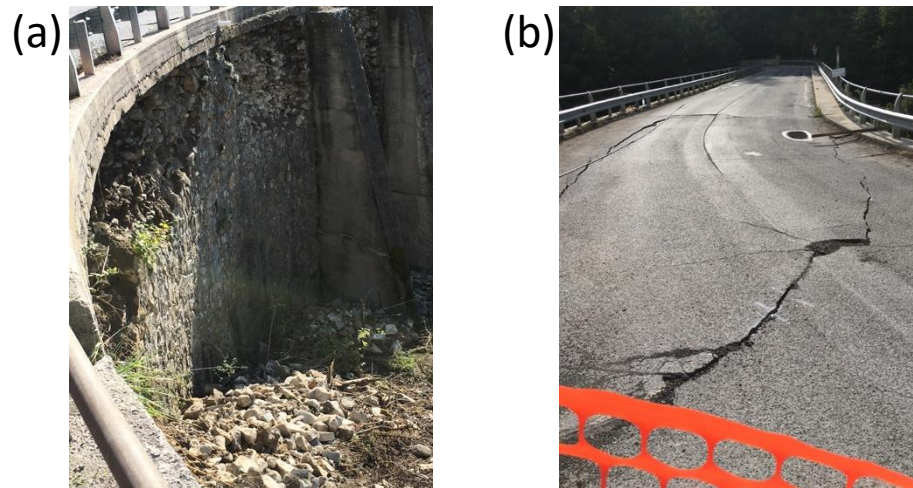
- RC bridge located along the Strada Provinciale (SP) 7 (km 16+150), in Boscomartese;
- RC bridge located along the SP 173, in the belt of Offida;
- RC bridge located along the SP 20, in Colle.
- Composite steel and concrete bridge located along the SP in Colle, also known as Ponte Ramazzotti.

We found no evidence of damage to any of the contemporary bridges. We found damage to two arch masonry bridges at the locations shown in Figure 1. One of these bridges is near the village of Tufo on the Roman-era Trisungo route. The bridge had some cracking and spalling of exterior masonry elements, apparently exacerbated by settlement at the east support of the arch, as shown in Figure 7.



**Figure 7.** Roman-era bridge along the Trisungo route (Tufo area – Arquata del Tronto): (a) spalling and (b) cracking of exterior masonry elements. (N42.735389, E13.253611, September 7 2016).

A second damaged bridge is the Ponte a Tre Occhi (Three eyes) near Amatrice, which is a critical lifeline for the access to the village of Amatrice. One of the arches exhibited cracking, but the most severe damage occurred to approach structures of *muratura a sacco* masonry construction, in which portions of the outer layer of masonry were lost, leading to lateral relaxation and settlement of the bridge deck (Figure 8). The bridge remained closed as of the date of our last site visit (8 September 2016).



**Figure 8.** Ponte a Tre Occhi (Amatrice): (a) spalling of the outer layer of the masonry and (b) settlement of the road surface. (N42.62067, E13.290278, September 8 2016).

### Within-Village Damage to Building Structures

The earthquake produced devastating effects on dwellings in the villages of Arquata del Tronto, Accumoli, Amatrice, and Pescara del Tronto, the overwhelming majority of which are of masonry construction. The motivation of the GEER work in these, and other, areas was to document patterns of damage and non-damage, and specific case histories of special interest. From a geotechnical and seismological perspective, these findings have potential application for inference of ground motion spatial distribution, including site effects (especially topographic effects). Moreover, as some of the initial engineering researchers to access the region, we also recognized our responsibility to support broader reconnaissance efforts for this earthquake in relation to the performance of buildings.

Prior to field deployment of the GEER team, we overlaid locations of structural damage as assessed by the Copernicus EMS Rapid Mapping service (<http://emergency.copernicus.eu/>), which is based on high-resolution orthophotos taken on 08/25/2016 at 10 UTC, with a pixel separation distance of 0.1m. That mapping delineated for each structure a damage classification as follows:

- Red: destroyed;
- Light red: highly damaged;
- Orange: moderately damaged;
- Yellow: Negligible to slight damage.

This mapping was provided for 29 villages by the Copernicus project.

We secured access to the 35 villages shown in Figure 1, and performed detailed mapping at three of these villages: (Arquata del Tronto, Pescara del Tronto, and Tufo). One objective of the detailed

mapping was to validate the Copernicus damage assessments. Analysis of this data remains in progress.

We also performed detailed aerial imaging of the Pescara del Tronto village, which will support the development of a 3D model for subsequent damage studies. Aerial flights were not authorized over the structures in the village of Accumoli during our reconnaissance. Close range photogrammetric survey for Digital Surface Model generation was also completed for the Sant'Agostino church and its bell tower in Amatrice (Figure 9), which can be used to generate 2D or 3D models of the structure.



**Figure 9.** 3D textured model of the Sant'Agostino church and its bell tower in Amatrice. (N42.628, E13.2914, September 7 2016, Figure prepared by: F Chiabrando, V Di Pietra, and N Grasso).

A consistent pattern of our observations pertains to the effectiveness of retrofitting for collapse avoidance of masonry structures. In the most heavily impacted villages, retrofitting was sparse. One compelling example is to compare Amatrice with Norcia. The devastation in Amatrice is well



known, and retrofit in the historic center was lacking. The ground motions in Norcia appear to have been of similar or even greater intensity (Figure 4), but damage was sparse. Most masonry dwellings in Norcia have been retrofitted under a government program implemented following earthquakes in 1979 and 1997.

We observed a number of retaining wall failures in the villages of Amatrice, Accumoli, and Pescara del Tronto. These will be further documented in our Ver 2 report.

## Next Steps

We have passed along our data resources and findings to the EERI team, which mobilized to the area 12-16 September 2016 in collaboration with Reluis and the EUcentre. We will maintain contact as we continue work on our Ver 2 report.

## Acknowledgements

The work of the GEER Association, in general, is based upon work supported in part by the National Science Foundation through the Geotechnical Engineering Program under Grant No. CMMI-1266418. Any opinions, findings, and conclusions or recommendations expressed in this material are those of the authors and do not necessarily reflect the views of the NSF. The GEER Association is made possible by the vision and support of the NSF Geotechnical Engineering Program Directors: Dr. Richard Frigaszy and the late Dr. Cliff Astill. GEER members also donate their time, talent, and resources to collect time-sensitive field observations of the effects of extreme events.

Sponsorship of GEER activities was also provided by the B. John Garrick Institute for the Risk Sciences at UCLA and the NSF I/UCRC Center for Unmanned Aircraft Systems (C-UAS) at BYU under Project BYU13-03.

Sponsorship of the Italian Geotechnical Society was also provided. Support from the Department of Civil Protection (Engs. Paola Pagliara and Paola Bertuccioli) is also kindly acknowledged.

The technical support by Consortium ReLUIIS (Network of Italian Laboratories of Earthquake Engineering), headquartered in the University of Naples Federico II, Italy, and the Department of Engineering of University of Sannio, Benevento, Italy, is also acknowledged.

We are grateful to Dr. Tadahiro Kishida at the Pacific Earthquake Engineering Research Center (PEER) for downloading and processing ground motions from the event sequence. These data were useful for planning reconnaissance activities and for the preparation of this report.

Tony Fierro and Luciano Mignelli, current master's students at University of Molise are acknowledged for their contributions to surface fault rupture data collection and synthesis. We thank Maria Luisa Bello and Maria Elena Sanseverino, master's students at University of Sannio for their contribution to damage to building structures data collection.

Support for one team member was provided by EUcentre (European Centre for Training and Research in Earthquake Engineering; Pavia, Italy).

## References

Ancheta T.D., Darragh R.B., Stewart J.P., Seyhan E., Silva W.J., Chiou B.S.-J., Wooddell K.E., Graves R.W., Kottke A.R., Boore D.M., Kishida T., and Donahue J.L. (2014). NGA-West2 database. *Earthquake Spectra*, **30**, 989-1005.

CERI working group: S. Martino, P. Caporossi, M. Della Seta, C. Esposito, A. Fantini, M. Fiorucci, R. Iannucci, G.M. Marmoni, P. Mazzanti, S. Moretto, S. Rivellino, R.W. Romeo, P. Sarandrea, F. Troiani, C. Varone (2016). Sisma Amatrice Effetti- Interazione rete infrastrutturale. Available at <http://www.ceri.uniroma1.it/>

Galadini F., and Galli P. (2003). Paleoseismology of silent faults in the Central Apennines (Italy): the Mt. Vettore and Laga Mts. faults. *Annals of geophysics*, **46**(5) 815-836.

Gruppo di Lavoro INGV sul terremoto di Amatrice (2016). Primo rapporto di sintesi sul Terremoto di Amatrice Ml 6.0 del 24 Agosto 2016 (Italia Centrale), doi: 10.5281/zenodo.61121 (in italian)

ISPRA (2016). Report attività svolta da ISPRA in data 25-26/08/2016, A (in Italian).available at <http://www.isprambiente.gov.it/files/notizie-ispra/notizie-2016/sisma-italia-centrale/ReportattivitavoltadaISPRAindata26.pdf>.

Pizzi A., and Galadini F. (2009). Pre-existing cross-structures and active fault segmentation in the northern-central Apennines (Italy). *Tectonophysics* **476** (1), 304-319.